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The effect of horizontal resolution on the representation of the global monsoon annual cycle in Atmospheric General Circulation Models

Lixia Zhang^{1,2}, Tianjun Zhou¹, Nicholas P. Klingaman³, Peili Wu⁴, Malcolm Roberts⁴

¹*LASG, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing, China*

²*Collaborative Innovation Center on Forecast and Evaluation of Meteorological Disasters, Nanjing University of Information Science & Technology, Nanjing, China*

³*National Centre for Atmospheric Science and Department of Meteorology, University of Reading, Reading, United Kingdom*

⁴*Met Office Hadley Centre, Exeter, United Kingdom*

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Corresponding author:

Dr. Lixia Zhang

LASG, Institute of Atmospheric Physics, Chinese Academy of Sciences,
Beijing 100029, China.

Phone: 86-10-8299-5452

Fax: 86-10-8299-5172

Email: lixiazhang@mail.iap.ac.cn

Abstract

The sensitivity of the representation of the global monsoon annual cycle to horizontal resolution is compared in three Atmospheric General Circulation Models (AGCMs): the Met Office Unified Model-Global Atmosphere 3.0 (MetUM-GA3), the Meteorological Research Institute AGCM3 (MRI-AGCM3) and Global High Resolution AGCM from the Geophysical Fluid Dynamics Laboratory (GFDL-HiRAM). For each model, we use two horizontal resolution configurations for the period 1998–2008. Increasing resolution consistently improves simulated precipitation and low-level circulation of the annual mean and the first two annual cycle modes, as measured by pattern correlation coefficient and Equitable Threat Score. Improvements in simulating the summer monsoon onset and withdrawal are region-dependent. No consistent response to resolution is found in simulating summer monsoon retreat. Regionally, increased resolution reduces the positive bias in simulated annual mean precipitation, the two annual-cycle modes over the West African monsoon and Northwestern Pacific monsoon. An overestimation of the solstitial mode and an underestimation of the equinoctial asymmetric mode of the East Asian monsoon are reduced in all high-resolution configurations. Systematic errors exist in lower-resolution models for simulating the onset and withdrawal of the summer monsoon. Higher resolution models consistently improve the early summer monsoon onset over East Asia and West Africa, but substantial differences exist in the responses over Indian monsoon region, where biases differ across the three low-resolution AGCMs. This study demonstrates the importance of a multi-model comparison when examining the added value of resolution and the importance of model physical parameterizations for the Indian monsoon simulation.

Key Words : global monsoon, high resolution modeling, annual cycle, AMIP

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1. Introduction

Monsoons are characterized by a seasonal wet-dry contrast and a reversal of prevailing winds, due to the reversal of land-sea and inter-hemispheric thermal contrasts forced by the annual cycle of solar heating. The global monsoon is a dominant mode of annual variability of the global tropical circulation and is a response of the coupled climate system to annual variations in solar forcing (Wang and Ding 2008). Global monsoon regions sustain nearly two-thirds of the world’s population and support some of the largest and fastest growing urban areas. The simulation and prediction of the annual cycle of monsoon circulation and precipitation is critically important, because of the effect of monsoons on hydrology, agriculture and local livelihoods and economies.

Although global monsoon regions are connected through mass conservation (Trenberth et al. 2000), each regional monsoon has unique characteristics due to the local interactions among the land surface, ocean, convection and synoptic weather systems, especially over Asian-Australian monsoon region. The well-recognized Asian summer monsoon system typically first onsets over the Bay of Bengal (BoB) in early May, followed by the South China Sea in mid-May and then India in early June (Wu and Zhang 1998; Xu and Chan 2001; Mao and Wu 2007). The onset over the BoB is often preceded by the development of a monsoon onset vortex (Krishnamurti and Ramanathan 1982; Wu et al. 2012). A numerical study by Liu et al. (2002) demonstrated that the onset of South China Sea monsoon is due to a Rossby wave train triggered by the strong latent heating during the BoB monsoon onset. The seasonal march of the East Asian summer monsoon (EASM) displays a step-wise northward and northeastward advance. From early May to mid-May, the ridge line of the western Pacific subtropical high (WPSH) is located along 15°N, and southern China experiences a pre-monsoon rainy season. Later, the WPSH exhibits two northward jumps in June and July with the ridge line located at 20°N and 25°N, respectively, the monsoon

rain band extends abruptly from the Indochina Peninsula – the South China Sea – from the Philippines to the Yangtze River valley in early to mid-June, and the Meiyu (or Baiu in Japan and Changma in Korea) begins. The monsoon penetrates into northern China (34–41°N) in mid-July, and the monsoon rainy season there lasts for one month and ends in early-middle August (Zhou et al. 2009). For the Australian summer monsoon onset in December, there are four major contributing factors, including land-sea thermal contrast, barotropic instability, arrival of the Madden-Julian oscillation (MJO) and intrusion of a mid-latitude trough (Hung and Yanai 2004). The mean onset of the summer monsoon over West Africa is 24th June, which is linked to an abrupt latitudinal shift of the Inter-Tropical Convergence Zone (ITCZ) from a quasi-stationary location at 5°N in May-June to another quasi-stationary location at 10°N in July-August (Sultan and Janicot 2003). For American summer monsoon, the northward rainbelt movement over southwestern North America from mid-to-late June is associated with the northward progression of the ITCZ (Yu and Wallace 2000; Higgins and Shi 2001), while the onset of South American monsoon is related to eastward displacement of the South Atlantic subtropical high; it is also affected by intraseasonal variability through low-frequency trough or ridge occurrence over southern Brazil (Raia et al. 2008). The unique features of regional monsoons and their onset have been rigorous metrics for gauging climate models' performances.

The complexity of monsoon systems presents great challenges to simulate the climatological seasonal means and annual cycles of the monsoon (Sperber et al. 2013; Zhou et al. 2017), although substantial efforts have been made to improve model physics and dynamics in the past several decades. Model biases in the simulation of the monsoons limit the fidelity of monsoon predictions and projections (Dong et al. 2016). Sperber et al. (2013) evaluated the performance of 25 coupled climate models from Coupled Model Intercomparison Project Phase 5 (CMIP5) and 22 models from CMIP Phase 3 (CMIP3) in simulat-

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3 99 ing the Asian summer monsoon onset and retreat. The authors found an overall delayed
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5 100 onset over India in models, even though CMIP5 models showed improved fidelity relative
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7 101 to CMIP3. The onset and retreat definition in Sperber et al. (2013) was based on rainfall
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9 102 thresholds; monsoon onset might not occur in models, or may be delayed relative to obser-
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11 103 vations, because of systematic dry biases. To avoid this issue, Sperber and Annamalai
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13 104 (2014) proposed a fractional accumulation method to evaluate the monsoon annual cycle;
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15 105 this method can be applied to many monsoon domains. The authors also found some sys-
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17 106 tematic errors in the phase of rainfall annual cycle: coupled climate models in CMIP5 have
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19 107 delayed onsets of summer rainfall over India, the Gulf of Guinea, and South America, but
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21 108 early onsets for the Sahel and the North America.
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27 109 Several studies have proposed techniques to improve monsoon simulation, including in-
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29 110 creasing horizontal resolution to capture more realistic large-scale atmospheric circulations
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31 111 and precipitation distributions (Hack et al. 2006; Roberts et al. 2009; Demory et al. 2014 ;
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33 112 Berckmans et al. 2013). The added value of resolution in GCMs has been widely verified
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35 113 for many aspects of monsoon simulation (Kitoh et al., 2008; Mizuta et al., 2012; Johnson et
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37 114 al., 2016). For example, an examination of version 5.1 of the National Center for Atmos-
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39 115 pheric Research Community Atmosphere Model (CAM5) at three resolutions showed a
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41 116 much better representation of the intensity-frequency structures of precipitation in steep-
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43 117 terrain regions over East Asia (Li et al. 2015). Higher resolution in CAM5 also improves
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45 118 the simulation of East Asian summer monsoon rain belt (Yao et al. 2017). Most investiga-
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47 119 tions of the benefits of increased resolution were based on one model, however the sensi-
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49 120 tivity to resolution may differ among models. Ogata et al. (2017) showed that the MRI-
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51 121 AGCM3 and the MetUM-GA3 produced less precipitation over the western Pacific with
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53 122 increasing resolution, but their sensitivity of Indian Ocean precipitation to resolution dif-
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55 123 fered. Zhang et al. (2017) compared three AGCMs, each with two resolution configura-
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tions, and showed the sensitivity of monsoon precipitation to resolution varied greatly among the models. To understand what aspects of resolution sensitivity are common among AGCMs, it is important to compare several AGCMs using the same metrics. Given the social and scientific importance of the onset and cessation of monsoon precipitation to local livelihoods and economies, it is desirable to find whether high resolution systematically improves the representation of monsoon onset and withdrawal. This is the major motivation of the current study.

The remainder of the paper is organized as follows. The model simulations, validation data and evaluation metrics are described in section 2. The improvement of high resolution in simulating the annual cycle modes and summer monsoon onset and withdrawal are assessed in section 3 and 4, respectively. Section 5 presents the summary and discussion.

2. Models, data and method description

2.1 Model simulations and observational validation datasets

We use daily precipitation and monthly wind data from simulations with three AGCMs, MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, each with two resolution configurations for the period 1998–2008. Details of the models and experiments can be found in Table 1. The model outputs of GFDL-HiRAM and MRI-AGCM3 were obtained from the Coupled Model Intercomparison Project phase 5 data archive (Taylor et al., 2012). The simulations of MetUM-GA3 are from the UPSCALE (UK on PRACE: weather-resolving Simulations of Climate for global Environmental risk) project (Mizielinski et al., 2014). There are very few different physical and dynamical settings in the MetUM-GA3 high-resolution configuration compared to its low-resolution counterparts, mostly related to numerical stability (Table 2 in Mizielinski et al. 2014). There are no physical parameter adjustments between the MRI-AGCM3-2H and MRI-AGCM3-2S (Endo et al. 2012; Kusunoki 2016). In GFDL-HiRAM, the shallow convective scheme and cloud microphysics are changed in

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C360 from C180 (Zhao et al. 2009; Chen and Lin 2012). We analyze the ensemble mean of all realizations of each model. Because of the large inter-model differences in dynamical cores and physics, it is hard to compare the resolution sensitivity across models. To exclude the impact of model dependence and focus more on resolution, we compared the high- and low-resolution configurations of the each model in this study.

The observed daily precipitation datasets used in this study include (1) Tropical Rainfall Measuring Mission (TRMM) 3B42 V7 (Huffman et al. 2007), to evaluate the climatological mean state of global monsoon precipitation, at $0.25^{\circ}\times0.25^{\circ}$ resolution; and (2) Global Precipitation Climatology Project V1.2 (GPCP), with a resolution $1.0^{\circ}\times1.0^{\circ}$ (Adler et al. 2003). In addition, monthly wind data from Climate Forecast System Reanalysis (CFSR; Saha et al. 2010) with a resolution of $0.5^{\circ}\times0.5^{\circ}$ and from National Centers for Environmental Prediction–National Center for Atmospheric Research (NCEP–NCAR) reanalysis with a resolution $2.5^{\circ}\times2.5^{\circ}$ (Kalnay et al. 1996) are used as the circulation datasets. All datasets are interpolated to the N216 resolution of MetUM-GA3 (approximately $0.83^{\circ}\text{longitude}\times0.55^{\circ}\text{latitude}$) using a distance-weighted interpolation method, to facilitate comparison. The longest common period of 1998–2008 covered by all simulations and observations is selected in the following analysis.

2.2 Evaluation metrics on global monsoon annual cycles

In this study, two aspects of the annual cycle of the global monsoon are evaluated: (1) the mean climate and annual cycle modes and (2) monsoon onset and withdraw. As proposed by Wang and Ding (2008), the annual mean, the solstitial (symmetric) and equinoctial (asymmetric) modes of the annual cycle, and the global monsoon domain can be used as metrics to gauge model performance for simulating the mean climate and annual cycle. The solstitial mode is represented by the differences in precipitation or circulation between

173 June–September and December–March means; the equinoctial asymmetric mode is repre-
 174 sented by the differences between April–May and October–November. The global monsoon
 175 domain is delineated both by the monsoon precipitation index (MPI) and zonal westerly
 176 wind index (MWI). MPI is defined as the annual range of precipitation normalized by its
 177 annual mean. MWI is similar, but using 850 hPa zonal wind (U850). The monsoon precipi-
 178 tation domains are the areas where the annual range of precipitation exceeds 300mm and
 179 MPI exceeds 0.5, whereas monsoon wind domains are defined as WMI exceeding 0.5
 180 without any restriction on the magnitude of annual range. The Annual range is the differ-
 181 ence between May–September (MJJAS) and November–March (NDJFM) means in the
 182 Northern Hemisphere, or NDJFM minus MJJAS in the Southern Hemisphere.

183 Previous studies have proposed several definitions of monsoon onset and withdrawal.
 184 The commonly used definition is threshold-based (Wang and Lin 2002). One of the weak-
 185 nesses of the threshold-based techniques is that monsoon onset might be delayed or accel-
 186 erated by dry or wet biases, respectively, in the model, even though the model may have a
 187 realistic annual cycle amplitude (Sperber et al. 2013; Sperber and Annamalai 2014). Mod-
 188 els with dry biases are especially at a disadvantage since the observed threshold for defin-
 189 ing monsoon may never be reached in such models, including MetUM-GA3 in this study
 190 (Bush et al. 2015; Johnson et al. 2016). Since the biases differ among the three models
 191 used here, it is important to find a relatively fair way to evaluate onset and withdrawal in
 192 models. Thus, a fractional accumulation approach is employed here. The monsoon onset
 193 and withdrawal are defined as the pentad when the fractional accumulation first be-
 194 comes ≥ 0.2 and ≥ 0.8 , respectively (Sperber and Annamalai 2014). The fractional accu-
 195 mulated precipitation in a given pentad is the accumulated rainfall at that pentad divided by
 196 the total accumulated rainfall at the end of the year. In this study, the pentads for the

Southern African, Australian and South American domains have been reordered to run from July 2 to June 27.

To quantitatively show the sensitivity of the simulated monsoon domain to resolution, the Equitable Threat Score (ETS, Jolliffe and Stephenson 2003) is used in this study to measure the simulation skill of the monsoon domain, which can be evaluated as a binary event.

$$ETS=(a-ar)/(a+b+c-ar) \tag{1}$$

where $ar=(a+b)(a+c)/n$, represents the skill of a random prediction. a, b, and c are the counts of hits, false alarms and misses of monsoon domain, respectively, and n is the number of model gridpoints in the region (45°S-45°N). The definition of each variable in equation (1) can be found in Table 2. ETS ranges from 0 to 1 with perfect score of 1.

3. Annual mean and annual cycle modes

To investigate whether the improvement from higher-resolution is consistent across the three models, we analyze the biases in the lower-resolution configuration and the differences between the high- and low- resolution configurations. The bias in simulated annual mean precipitation and improvements from higher resolution are presented in Fig.1. In observations (blue lines in left panel of Fig.1), substantial precipitation falls in the Indo-Pacific warm pool (60°E-180°), ITCZ and South Pacific Convergence Zone (SPCZ). In general, the observed centers are well captured by the low-resolution configuration of each AGCM, with an overall positive bias over most of the tropics (blue lines in right panel of Fig.1). In contrast, there are several common biases among the three low-resolution AGCMs, including overestimations over the southwest Indian Ocean, SPCZ, northern edge of the ITCZ, and Northwestern Pacific Ocean, with underestimations over the southeastern Indian Ocean (Fig.1 a, c, e). Unique biases in each AGCM are also found, particularly in

221 MetUM-GA3 N216, which underestimates Indian and Maritime Continent precipitation
 222 (Fig.1a). A positive bias over the western central Pacific (5°S - 5°N , 120°E - 180°) is seen in
 223 MRI-AGCM3-2H, while negative biases are present in MetUM-GA3 N216 and GFDL-
 224 HiRAM-C180. A consistent response to higher resolution among the three models is sup-
 225 pressed precipitation over the north of Maritime Continent (MC) and western Pacific (5°S -
 226 15°N , 120° - 160°E), which partly reduces the wet bias over the Northwestern Pacific (5° -
 227 15°N , 120° - 160°E) in the three low-resolution models. This consistent response to finer
 228 resolution also has been illustrated by Ogata et al. (2017), which compared MetUM-GA3
 229 and MRI-AGCM3. This is further verified with one additional model, GFDL-HiRAM. As
 230 suggested by Ogata et al. (2017) and Johnson et al. (2016), the reduced precipitation over
 231 the north of MC and western Pacific may be related to better-resolved steep topography
 232 over MC and Indochina, with moisture convergence and precipitation on the windward
 233 side of the orography, which leads to decreased moisture availability on the leeward side in
 234 summer and reduced precipitation over the north of MC and western Pacific. This can be
 235 partly seen from the difference between higher- and lower-resolution models in simulating
 236 the solstitial mode in Fig.2. Note that the low-resolution of both MetUM-GA3 and GFDL-
 237 HiRAM underestimate the precipitation over MC, and the consistent response to resolution
 238 makes this dry bias even worse.

239 As in Fig. 1, the biases in low-resolution models and improvement in higher- resolu-
 240 tion configurations for the solstitial mode are shown in Fig. 2. The observed spatial pattern
 241 shows an asymmetric contrast, with positive centers in Northern Hemisphere monsoon re-
 242 gions (Indian-Pacific-East Asia, West Africa and northeast tropical Pacific (5° - 20°N , 120° -
 243 60°W)) and negative centers in Southern Hemispheric monsoon regions (South Africa,
 244 Australia and South America; lines in left panel of Fig.2). The centers are well reproduced
 245 by all low-resolution models, but all show underestimations over East Asia, the northeast

246 tropical Pacific, South Africa and Australia region and overestimations over the northwest
 247 Pacific (shading in left panel of Fig.2). Higher resolution reduces this consistent simulation
 248 bias, particularly over the northwest Pacific and northeast tropical Pacific (shading in left
 249 panel of Fig.2). The improvement, however, is far smaller than the low-resolution model
 250 bias. The biases in the three low-resolution models differ greatly over the Indian monsoon
 251 region, where the solstitial mode is dramatically smaller than observation in MetUM-GA3
 252 N216, greater in MRI-AGCM3-2H, and close to observation in GFDL-HiRAM-C180. The
 253 sensitivity of the solstitial mode in each model to resolution also differs over the Indian
 254 monsoon region. The inconsistency of model biases and sensitivity to resolution indicate
 255 the strong role for sub-grid scale physical parameterizations for simulating the Indian mon-
 256 soon precipitation.

257 The biases in the simulated equinoctial asymmetric mode of the annual cycle and cor-
 258 responding improvements in high-resolution models are shown in Fig. 3. In observations,
 259 negative values are centered in the zonal belt along (10°N-20°N), while positive values are
 260 centered to its south (0°-10°N) and in Southeast China. The centers over the Pacific repre-
 261 sent the asymmetric location of the spring and fall ITCZ (Wang and Ding 2008). The posi-
 262 tive values over Southeast China shows the region's characteristic spring-persistence sea-
 263 son in April and May (Li et al. 2017). All low-resolution models capture the asymmetric
 264 pattern of spring and fall precipitation, but a systematic overestimation over the Northern
 265 Hemisphere is seen in all models, with negative biases north of the observed negative cen-
 266 ter and positive biases near the observed positive centers. Positive biases over West Africa
 267 are found in all low-resolution models. In high-resolution models, the overestimation over
 268 Northwestern Pacific, West Africa and Southeast China is suppressed in all models; consis-
 269 tent improvements are also found over the southern Maritime Continent. This indicates
 270 increased resolution improves simulated precipitation in transition seasons.

271 To quantify the improvement in high-resolution models, the pattern correlation coef-
 272 ficient (PCC) and root-mean-square-error (RMSE) of simulated precipitation over (45°S-
 273 45°N) against TRMM are shown in Fig.4 a,d,g. The simulated 850 hPa winds for the an-
 274 nual mean and the first annual cycle modes against CFSR reanalysis are also presented in
 275 Fig.4. We compare the metrics for GPCP against TRMM and for NCEP-NCAR against
 276 CFSR to show the observational uncertainty. In general, all low-resolution models simu-
 277 late well the observed precipitation and 850 hPa wind patterns. For example, the PCCs
 278 (RMSEs) of simulated annual mean precipitation and 850 hPa zonal (U850) and meridio-
 279 nal wind (V850) in low resolution models (blue markers in Fig.4a-c) range from 0.85 to
 280 0.90 (1.2 to 1.6 mm/day), 0.97 to 0.99 (0.75 to 1.2 m/s) and 0.88 to 0.94 (0.47 to 0.60 m/s),
 281 respectively, which are close to those PCCs for GPCP and NCEP-NCAR against TRMM
 282 and CFSR, respectively. The PCCs increase consistently in all three high-resolution models
 283 (red markers) compared with their low-resolution counterparts (blue markers) for the three
 284 annual-cycle metrics. However, the model bias as demonstrated by RMSE appears insensi-
 285 tive to increasing resolution. The PCCs of annual precipitation, U850, V850 simulated by
 286 high-resolution models are systematically shifted to right of the low-resolution ones, in-
 287 creasing to 0.95~0.97, nearly 1.0 and 0.96~0.98, respectively. The PCCs in high-resolution
 288 models are even higher than PCCs of GPCP/NCAR, although the RMSEs in these
 289 simulations are far larger than the reanalysis. The same results are also found in the first
 290 two leading modes. High resolution aids in representing details of precipitation and circula-
 291 tion distributions, but not in reducing model biases.

292 Following the global monsoon domain definition proposed by Wang and Ding (2008),
 293 the distributions of precipitation annual range and monsoon domain based on precipitation
 294 are obtained and shown in Figure 5. In TRMM, the monsoon systems are well depicted by
 295 this definition, including the Asian–Australian monsoon, the North and South African

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3 296 monsoons, and the North and South American monsoons (Fig.5a). The domain obtained
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5 297 from GPCP is similar to TRMM except at the edges, particularly over the northern edge of
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7 298 the northwestern Pacific monsoon region and the southern edge of the southwestern Indian
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9 299 Ocean monsoon region, where the domain areas are relatively larger in GPCP than in
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11 300 TRMM. The North American monsoon region (120°W - 60°W , 0° - 20°N) in all low-
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13 301 resolution models is smaller and the western boundary is about 30 degrees eastward shifted
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15 302 relative to observations, while the simulated northwestern Pacific monsoon is wider and
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17 303 stronger than observations (left column in Fig.5). With increased resolution, the northwest-
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19 304 ern Pacific monsoon region shrinks relative to the low-resolution models, particularly for
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21 305 MRI-AGCM3, which extends to 170°W in MRI-AGCM3-2H but only to 170°E in MRI-
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23 306 AGCM3-2S. Each model has unique biases, such as the absence of the Indian monsoon
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25 307 region in MetUM-GA3 N216 due to its dry bias in India, a westward extended West Afri-
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27 308 can monsoon (60°W - 60°E , 0° - 20°N) in MetUM-GA3 N216 and MRI-AGCM3-2H due to a
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29 309 wider and stronger ITCZ over the Atlantic Ocean, and an unrealistic monsoon region over
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31 310 the southern tropical Atlantic Ocean in MRI-AGCM3-2H and GFDL-HiRAM-C180.
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33 311 Those biases remain in their high-resolution counterparts, indicating little influence of res-
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35 312 olution increase. Similar analysis of the global monsoon domain based on U850 revealed
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37 313 no obvious improvements from increased resolution (Figure not shown).
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45 314 To quantitatively show the sensitivity of the simulated monsoon domain to resolution,
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47 315 the Equitable Threat Score (ETS, Jolliffe and Stephenson 2003) for global and individual
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49 316 regional monsoon domains was computed, using precipitation and U850 (Figure 6). Gener-
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51 317 ally, higher-resolution configurations show slightly increased ETS for global monsoon pre-
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53 318 cipitation (wind): from 0.53 to 0.54 (0.63 to 0.67) in MetUM-GA3, 0.56 to 0.57 (0.68 to
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55 319 0.69) in GFDL-HiRAM and 0.58 to 0.62 (0.73 to 0.77) in MRI-AGCM3 models, respec-
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57 320 tively. The ETS of the three high-resolution models for regional monsoon precipitation
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domains are not always higher than their low-resolution counterparts. Consistent improvements with resolution are shown only for the South African monsoon domain, with the largest increase from 0.59 to 0.64 in MetUM-GA3. The ETS for regional monsoon wind domains are all higher in the high-resolution configurations. Note that the ETS metrics of U850 and precipitation are not correlated. The skill scores of U850 are higher than those of precipitation for both global and regional monsoons. Even though the higher-resolution models show increased ETS for U850 for all regional monsoons except the East Asian monsoon, the ETS metrics for precipitation do not increase (e.g., for the Indian monsoon).

4. Monsoon onset and withdrawal

Cumulative rainfall, which is the sum of rainfall up to a given pentad, averaged over each monsoon region is presented in Fig.7 to show the temporal characteristics of the rainfall and the model bias. A rapid increase in cumulative occurs in all monsoon regions. However, the initial pentad for the rapid increase differs among regions, indicating different onset times in each region. For the Asian-Australian monsoon system, rapid precipitation increase starts around pentad 30 in India, pentad 25 in Southeast Asia and East Asia, and pentad 10 in Australia. Over the African and American monsoon systems, an onset is observed in the pentad 30 in West Africa and North America, and pentad 5 in South Africa and South America.

An overall overestimation of annual total precipitation is presented over all regional monsoons except for India and Southeast Asia. For Indian monsoon, there is a large disparity among the models in simulated total rainfall. There is an obvious wet bias (50% wetter than TRMM) in MRI-AGCM3, a dry bias (50% less than TRMM) in MetUM-GA3 AGCMs and a relatively small bias in GFDL-HiRAM (Fig. 7a). The biases over Southeast Asia are similar to those over India, but with smaller magnitudes (Fig. 7b). The evolution of precipitation accumulation shows that the bias is relatively small in the beginning of

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year, but starts to increase around the monsoon onset pentad.

The observed and simulated spatial distributions of monsoon onset are shown in Fig.8. The monsoon onset is the pentad when fractional accumulation first becomes ≥ 0.2 (Sperber and Annamalai 2014). We also examined the distributions of monsoon onset defined by Wang and Lin (2002), in which onset is determined as the first pentad with the relative rainfall rate to January mean in NH and July in SH exceeds 5 mm d^{-1} . This threshold is never reached in MetUM-GA3 over India because of the systematic dry bias (Fig.1, Fig. 2 and Fig.7a). However, India is defined as a monsoon region in MetUM-GA3 based on the monsoon wind index (Figure not shown). Thus, it is reasonable to use the fractional accumulation to define monsoon onset and withdrawal. To focus on monsoon regions, only the distributions in the observed monsoon regions based on the monsoon precipitation index are shown here. In observations (Fig.8a-b), the onset pentad shows a gradual increase from the equator to poleward, indicating later onset in the higher-latitude monsoon regions in both hemispheres. In the NH, the earliest onset is pentad 25 in early May over the southern edge of the Northern Hemispheric monsoon region and South China, while in the SH the earliest onset is pentad 68 in early December over the northern edge of Southwest Indian Ocean. The latest NH monsoon onset is at pentad 47 in mid-August over the northern edge of the North American and West African monsoons (pentad 15 in mid-March over the south edge of Australian monsoon in SH).

All models simulate well the spatial distribution of global summer monsoon onset. However, all low-resolution models simulate earlier monsoon onsets than the two observed datasets, such as over the East Asian, West Africa monsoon, North American and South American monsoon. Specifically, the area with onset by pentad 25 (hereafter P25 area) over the West African, North American and East Asian monsoons is broader than in the observation. In the high-resolution models, the P25 area slightly shrinks over the East

371 Asian and Western African monsoons, particularly apparent in MetUM-GA3. This differ-
 372 ence is also seen in the observations, in which the P25 area in GPCP is larger than in
 373 TRMM, potentially due to the different resolutions of two observed datasets. It indicates
 374 that high resolution improves the spatial distribution of precipitation and the delineation of
 375 the monsoon onset boundaries. In other monsoon regions, there are no clear advantages of
 376 increased resolution.

377 As for the global monsoon onset, we compute distributions of monsoon withdrawal,
 378 the pentad with fractional accumulation first becomes ≥ 0.8 , in simulations and observa-
 379 tions (Fig. 9). The observed monsoon withdraws from higher latitudes toward the equatori-
 380 al monsoon regions. In NH, the earliest withdrawal over East Asian monsoon is witnessed
 381 at pentad 48 in late August; the latest is over the southern edge of the Northern Hemi-
 382 spheric monsoon region at pentad 60 in late October. In the SH, the earliest is over South
 383 Africa at pentad 14 in mid-March and the latest over northern edge of Australia at pentad
 384 30 in late June. The observed monsoon withdrawal patterns are well captured by all mod-
 385 els, but with relatively later withdrawals over the Australian and North American mon-
 386 soons by two pentads. There are no obvious differences between the low- and high-
 387 resolution configurations, or between GPCP and TRMM. These results suggest monsoon
 388 withdrawal is less sensitive to resolution than monsoon onset.

389 To quantitatively identify the improvements from higher resolution, the summer mon-
 390 soon onset and withdrawal pentads averaged over each regional monsoon are shown in
 391 Figure 10. In observations, TRMM shows (red dots in Fig. 10) that the monsoon onsets
 392 (withdraws) at pentad 33 (52) over India, 30.5 (54) over Southeast Asia, 27.5 (49.5) over
 393 East Asia, 7 (28.5) over Australia, 32.8 (53.5) over West Africa, 8.2 (27) over South Afri-
 394 ca, 32.4 (56) over North America and 3.2 (27.2) over South America, respectively. Com-
 395 pared with TRMM, the onset pentads of all regional monsoons except the Indian monsoon

are systematically earlier in GPCP, the lower-resolution observational dataset, with the largest difference of up to two pentads earlier in the West Africa monsoon. However, the monsoon withdrawal dates from TRMM and GPCP are similar, with a maximum difference of 0.5 pentad. The difference between TRMM and GPCP may suggest that resolution has greater impact on monsoon onset than on withdrawal. For all domains the model dispersion far exceeds the differences between TRMM and GPCP. Here, we use the onset and withdrawal derived from TRMM as a reference due to its higher resolution than GPCP. All models show earlier onsets over East Asia, West Africa, South Africa and North America. The results are consistent with Fig.8 and Fig.9. Specifically, the largest bias is about 4.5 pentads in MRI-AGCM3-2H over East Asia, 4 pentads in GFDL-HiRAM-C180 over West Africa, 2 pentads in GFDL-HiRAM-C360 over South Africa and 4.5 pentads in GFDL-hIRAM-C360 over North America, respectively. Increased resolution consistently delays the monsoon onset over East Asia and West Africa in all AGCMs, matching well with the differences between TRMM and GPCP. However, these improvements from high resolution are only one pentad at most, far less than the biases of the corresponding low-resolution models (≤ 4.5 pentad). A consistently slightly advanced onset in higher-resolution models (1.5 pentad earlier at maximum over Southeast Asia) relative to their lower-resolution counterparts is shown over Southeast Asia, Australia, South Africa, North America and South America. For the Indian monsoon, both the biases of the lower-resolution configurations and the improvement in the high-resolution configurations differ greatly among the three AGCMs, indicating less resolution sensitivity of Indian monsoon onset.

5. Summary and discussion

In this study, the resolution sensitivity of the annual cycle of the global monsoon is investigated in three AGCMs, MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM. We com-

pared output from MetUM-GA3 at N216 (90-km) and N512 (40-km), MRI-AGCM3-2H (60-km) and MRI-AGCM3-2S (20-km), and GFDL-HiRAM-C180 (65-km) and GFDL-HiRAM-C360 (32-km). We found consistent and inconsistent responses across the three AGCMs to increasing resolution.

A summary plot about the model responses to horizontal resolution is presented in Fig.11. All low-resolution AGCMs are biased toward higher annual mean precipitation and stronger solstitial and equinoctial asymmetric modes. With increasing resolution, all AGCMs show consistent improvements in simulating the precipitation and low-level circulation of the annual mean and the first two annual cycle modes, as determined from the pattern correlation coefficient and ETS as denoted by red boxes in Fig.11. Regionally, higher pattern correlation coefficients are found in three high-resolution AGCMs for precipitation for the annual mean (ANN), solstitial mode (AC1) and the equinoctial asymmetric mode (AC2) for all regional monsoons, except the Australian (Aus) and North American monsoons (NA_m). The inconsistent response to increased resolution across the three AGCMs over North American and Australian monsoons may be partly due to their small monsoon area.

We found that improvements in simulating the summer monsoon onset and withdrawal are regionally dependent. No consistent response to resolution is found in simulating monsoon withdrawal. A consistent improvement in simulating East Asian (EAs), Southeast Asian (SAs) and Western African (Waf) monsoon onset are shown across three AGCMs, by delaying the monsoon onset over EAs and Waf and advancing over SAs. A consistently weakened equinoctial asymmetric mode over EAs and Waf with increasing resolution is seen across the three models, reducing wet biases in the transition seasons. Thus, the summer monsoon onset pentad of the two regional monsoons is systematically delayed in all high-resolution models relative to their low-resolution counterparts. In contrast, the im-

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provements of monsoon onset over SAs are consistently advanced across three models. Systematic errors and responses to resolution differ greatly across the three AGCMs for the simulated Indian (India) North American (NA_m) and South hemispheric regional monsoons (Aus, SA_f, SA_m), and the response to resolution differs across the three AGCM, indicating the dependence of these regions on physical parameterizations. Nevertheless, better representation on the Indian are shown in all three AGCMs.

Comparing the impact of resolution on the simulated regional monsoon precipitation annual cycle, this study demonstrates the importance of resolution for the East Asian monsoon and West African monsoon where improved annual cycles are simulated in the three AGCMs. The positive bias of AC2 over the West African monsoon region is caused by overestimated precipitation in spring; in CMIP5 models, this bias has been linked to the representation of African Easterly Waves (see Fig.6 in Martin and Thorncroft 2015). Both the AMIP and historical simulations of CMIP5 show wet biases over northern Africa in spring; this bias is significantly reduced in high-resolution models. A possible reason is that the finer resolution weakens the excessive African Easterly wave activity over northern Africa and thus reduces rainfall (Martin and Thorncroft 2015). Over the East Asian monsoon region, the summer monsoon rainfall, particularly the Meiyu-Baiyu rainbelt, is underestimated in many contemporary models (Kang 2004; Huang et al. 2013). Therefore, the simulated contribution of summer rainfall to the annual total is also underestimated, resulting in relatively faster fractional accumulations of precipitation in spring and earlier onset of EAs. As shown in Yao et al. (2017), as resolution increases, intensified northerly flow over the central northern areas of China and southerly flow to the south of Japan is found both from the AMIP simulations in CMIP5 and in CAM5 with different resolutions. Those changes with resolution are related to the topography driven barotropic Rossby waves downstream of the Tibetan Plateau (Yao et al. 2017). However, how the in-

creased resolution affects the regional circulation remains inconclusive and deserves further investigation.

Although several consistent responses to resolution are identified across the three AGCMs considered, large differences still exist, especially over the Indian monsoon region. Large differences in the sensitivity to resolution over the Indian monsoon region were found between MetUM-GA3 and MRI-AGCM3, particularly in synoptic and intra-seasonal variability, such as monsoon low-pressure systems (Ogata et al. 2017). Those inconsistencies across models demonstrate the importance of improving physical parameterizations to reduce these systematic errors. This paper highlights the need for further multi-model comparisons to determine the added value of horizontal resolution in climate simulations.

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Figure Captions:

Table 1 Model descriptions including model resolution, sea surface temperature (SST) forcing, realization numbers and selected simulation years.

Table 2. Schematic contingency table for computing the ETS metric for model simulations of the monsoon domain (after Table 3.1 in Jolliffe and Stephenson 2003)

Figure 1 Spatial distributions of the difference between low-resolution models and TRMM (units: mm day⁻¹, left column) and between high and low resolutions of the same model (units: mm day⁻¹, right column) in simulating annual mean precipitation. (a) MetUM-GA3 N216 minus TRMM, (b) MetUM-GA3 N512 minus MetUM-GA3 N216, (c) MRI-AGCM3-2H minus TRMM, (d) MRI-AGCM3-2S minus MRI-AGCM3-2H, (e) GFDL-HiRAM-C180 minus TRMM, (f) GFDL-HiRAM-C360 minus GFDL-HiRAM-C180. Blue lines in the left and right columns show the climatological mean precipitation from TRMM and the low-resolution models, respectively. The hatched area denotes where precipitation exceeds 7mm day⁻¹.

Figure 2 As in Figure 1, but for the spatial distribution of the solstitial mode (unit: mm day⁻¹), defined as the difference between June-September (JJAS) and December-March (DJFM) mean precipitation. The blue and red lines indicate positive and negative values, respectively.

Figure 3 As in Figure 1, but for the equinoctial asymmetric mode defined as the difference between April-May (AM) and October-November (ON) mean precipitation.

Figure 4 Scatter diagrams for the fidelity of simulated precipitation and 850hPa winds of (a)-(c) annual mean, (d)-(f) solstitial mode, and (g)-(i) equinoctial asymmetric mode against TRMM and CFSR. The abscissa and ordinate are the pattern correlation coefficients (PCC) and root-mean-square-error (RMSE), respectively. The fidelity of GPCP and NCEP2 against TRMM and CFSR are also shown in dots to represent the observa-

tional uncertainty. The stars, triangles and diamonds represent the results from MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution models in red and low-resolution in blue.

Figure 5 Climatological annual range of precipitation (shaded, units: mm day⁻¹) and global monsoon domain (contours) derived from (a) TRMM, (b) GPCP, (c) MetUM-GA3 N216, (d) MetUM-GA3 N512, (e) MRI-AGCM3-2H, (f) MRI-AGCM3-2S, (g) GFDL-HiRAM-C180, (h) GFDL-HiRAM-C360. The annual range is defined as the May to September (MJJAS) minus November to March (NDJFM) mean precipitation for the Northern Hemisphere, and NDJFM minus MJJAS for the Southern Hemisphere.

Figure 6. The ETS score of monsoon domain simulation over (a) global (GM), (b) Indian (India), (c) Southeast Asia (SAs), (d) East Asian (EAs), (e) Australian (Aus), (f) West African (Waf), (g) South African (SAf), (h) North American (NAf) and (i) South American (SAf) monsoon. The abscissa (ordinate) is ETS scores of simulated domain derived from precipitation (U850), respectively. The fidelity of GPCP against TRMM is shown in dots. The fidelity of GPCP and NCEP2 against TRMM and CFSR are also shown in dots to represent the observational uncertainty. The star, triangle and diamond represent the results from MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution models in red and low-resolution in blue.

Figure 7 Observed and simulated precipitation accumulations (unit: mm) from the first pentad to the end of year, and fractional accumulations (unit: %) around the monsoon onset pentad averaged over (a) Indian (India), (b) Southeast Asian (SAs), (c) East Asian (EAs), (d) Australian (Aus), (e) West African (Waf), (f) South African (SAf), (g) North American (NAf) and (h) South American (SAf) monsoon. Note: for Southern Hemisphere locations (Australian, South African, and South American monsoons) the pentads were reordered to July–June prior to the analysis. The black, red, blue and green

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lines represent the results of TRMM, MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution configurations in solid lines and low-resolution configurations in dashed lines.

Figure 8 Climatological monsoon onset pentad derived from (a) TRMM, (b) GPCP, (c) MetUM-GA3 N216, (d) MetUM-GA3 N512, (e) MRI-AGCM3-2H, (f) MRI-AGCM3-2S, (g) GFDL-HiRAM-C180, (h) GFDL-HiRAM-C360. The monsoon onset pentad is defined as the pentad when fractional accumulation first becomes ≥ 0.2 (Sperber and Annamalai 2014). Note: for Southern Hemisphere locations (Australian, South African, and South American monsoon) the pentads were reordered to July–June prior to the analysis.

Figure 9 As in Figure 8, but for the monsoon withdrawal pentad, defined as the pentad with fractional accumulation first becomes ≥ 0.8 (Sperber and Annamalai 2014). Note: for Southern Hemisphere locations (Australian, South African, and South American monsoon) the pentads were reordered to July–June prior to the analysis.

Figure 10 The monsoon onset (x axis) and withdraw pentad (y axis) averaged over (a) Indian (India), (b) Southeast Asian (SAs), (c) East Asian (EAs), (d) Australian (Aus), (e) West African (Waf), (f) South African (SAf), (g) North American (NAf) and (h) South American (SAf) monsoon. Note: for Southern Hemisphere locations (Australian, South African, and South American monsoon) the pentads were reordered to July–June prior to the analysis. The dots, stars, triangles and diamonds represent the results from the observations, MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution configurations in red and low-resolution configurations in blue.

Figure 11 Improvement of the simulation on the annual cycle of global monsoon precipitation with horizontal resolution. Red boxes indicate high-resolution consistently improves performance across all three AGCMs; blue boxes indicate improvements with

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3 681 resolution are inconsistent; orange boxes indicate a better simulation is shown in all
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5 682 three high-resolution AGCMs but the responses to resolution are inconsistent across
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10 684 low-resolution configurations of MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, re-
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12 685 spectively. The metric for the annual modes, including annual mean (ANN), solstitial
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19 688 are ETS and the onset/withdrawal pentad, respectively.
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Table 1. Model information description including model resolution, sea surface temperature (SST) forcing, realization numbers and selected simulation years.

Model	Resolution	SST	Realizations	Selected Years
MetUM-GA3	N216 (0.83°x0.55°)	OSTIA	3	1998-2008
	N512 (0.35°x0.23°)	OSTIA	5	1998-2008
GFDL-HiRAM	C180(0.625°x0.5°)	HadISST1	3	1998-2008
	C360(0.31°x0.25°)	HadISST1	2	1998-2008
MRI-AGCM3	2H(0.56°x0.56°)	HadISST1	2	1998-2008
	2S(0.18°x0.18°)	HadISST1	1	1998-2008

Table 2. Schematic contingency table for computing the ETS metric for model simulations of the monsoon domain (after Table 3.1 in Jolliffe and Stephenson 2003)

Event Observed			
Event Simulated	Monsoon domain	Non-monsoon domain	Total
Monsoon domain	a (Hits)	b (False alarms)	a+b
Non-monsoon domain	c (Misses)	d (Correct rejections)	c+d
Total	a+c	b+d	a+b+c+d=n

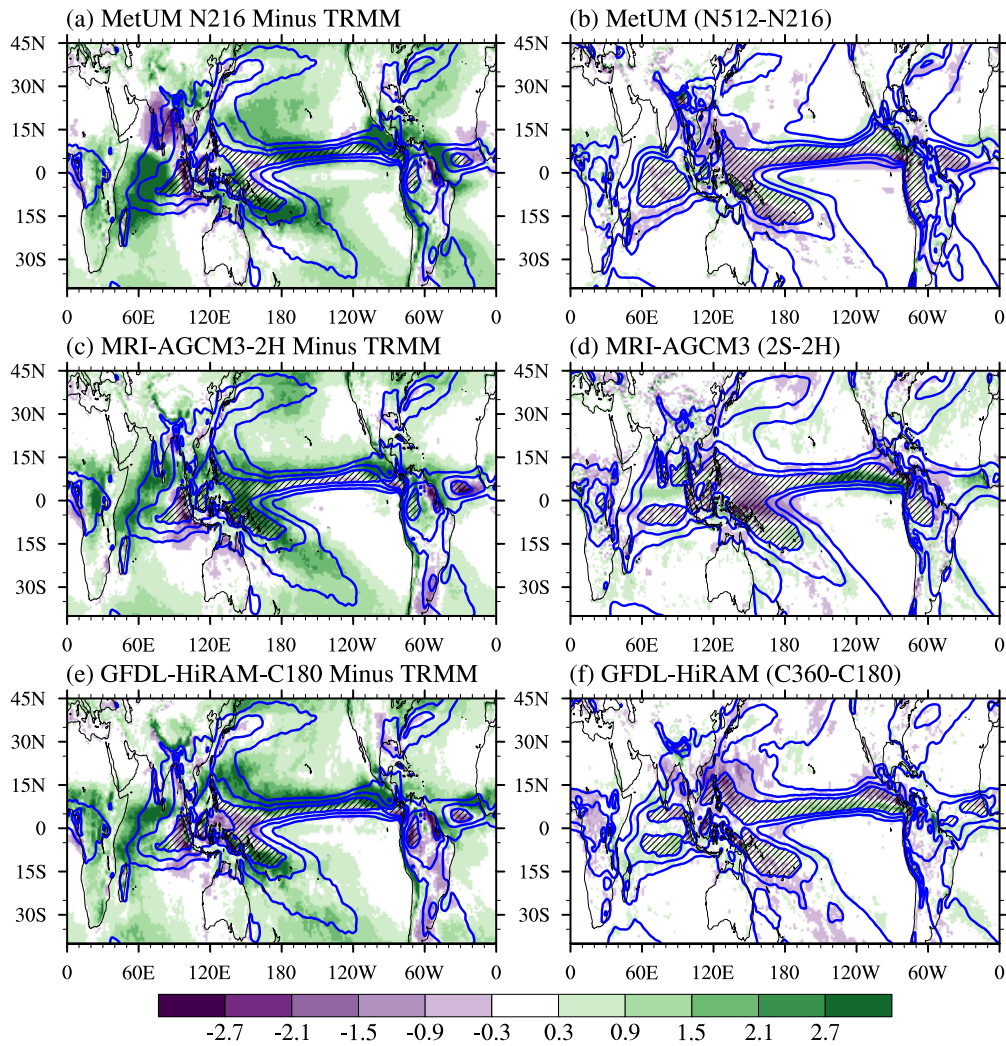


Figure 1 Spatial distributions of the difference between low-resolution models and TRMM (units: mm day⁻¹, left column) and between high and low resolutions of the same model (units: mm day⁻¹, right column) in simulating annual mean precipitation. (a) MetUM-GA3 N216 minus TRMM, (b) MetUM-GA3 N512 minus MetUM-GA3 N216, (c) MRI-AGCM3-2H minus TRMM, (d) MRI-AGCM3-2S minus MRI-AGCM3-2H, (e) GFDL-HiRAM-C180 minus TRMM, (f) GFDL-HiRAM-C360 minus GFDL-HiRAM-C180. Blue lines in the left and right columns show the climatological mean precipitation from TRMM and the low-resolution models, respectively. The hatched area denotes where precipitation exceeds 7mm day⁻¹.

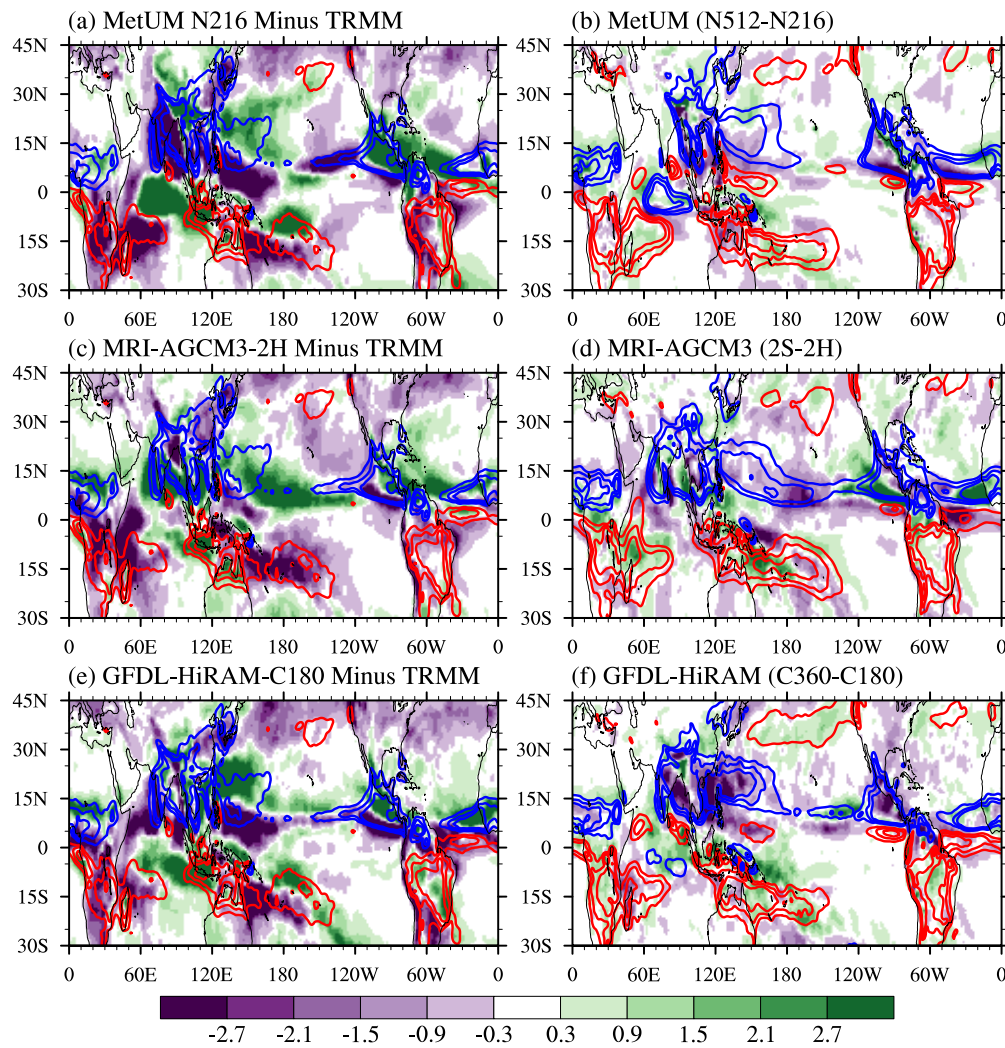


Figure 2 As in Figure 1, but for the spatial distribution of the solstitial mode (unit: mm day⁻¹), defined as the difference between June-September (JJAS) and December-March (DJFM) mean precipitation. The blue and red lines indicate positive and negative values, respectively.

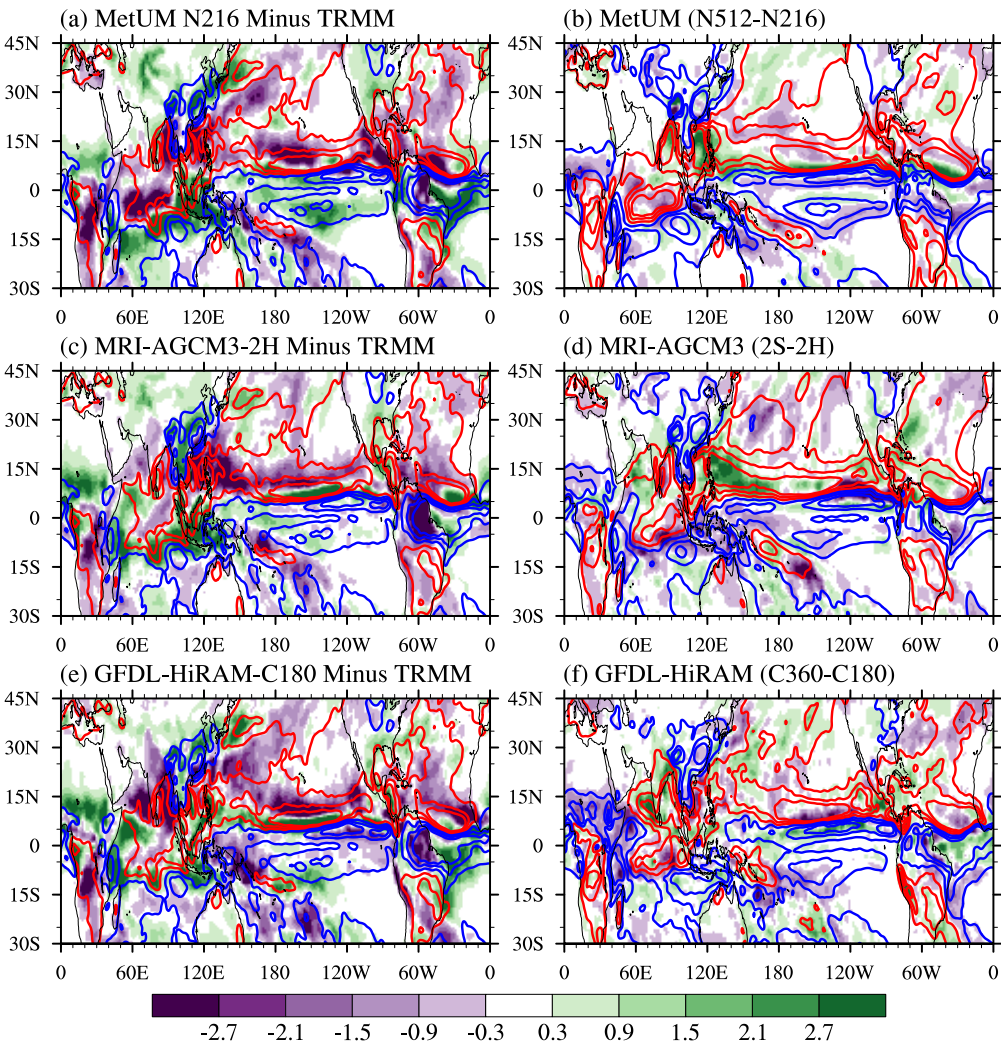


Figure 3. As in Figure 1, but for the equinoctial asymmetric mode defined as the difference between April-May (AM) and October-November (ON) mean precipitation.

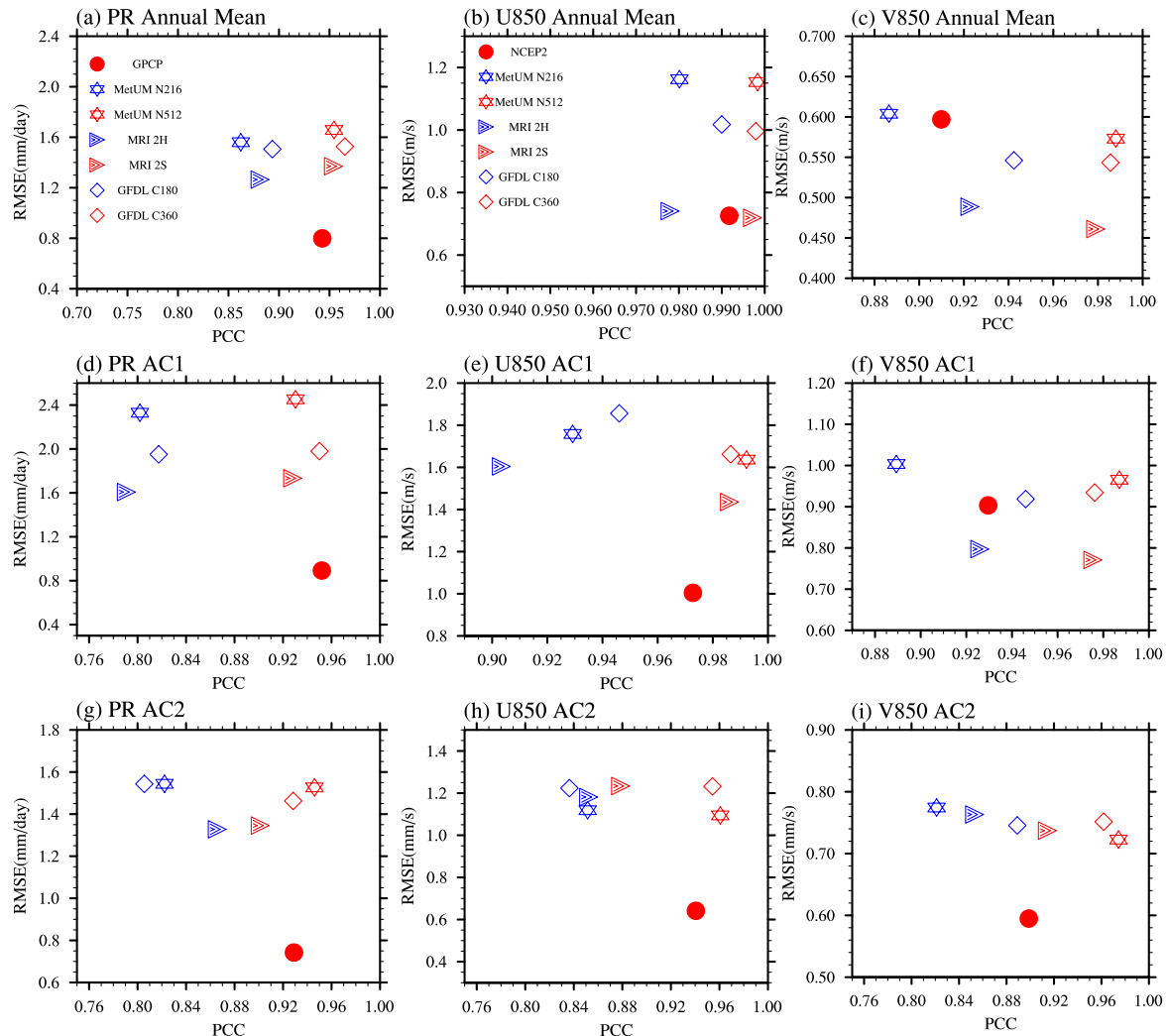


Figure 4 Scatter diagrams for the fidelity of simulated precipitation and 850hPa winds of (a)-(c) annual mean, (d)-(f) solstitial mode, and (g)-(i) equinoctial asymmetric mode against TRMM and CFSR. The abscissa and ordinate are the pattern correlation coefficients (PCC) and root-mean-square-error (RMSE), respectively. The fidelity of GPCP and NCEP2 against TRMM and CFSR are also shown in dots to represent the observational uncertainty. The stars, triangles and diamonds represent the results from MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution configurations in red and low-resolution configurations in blue.

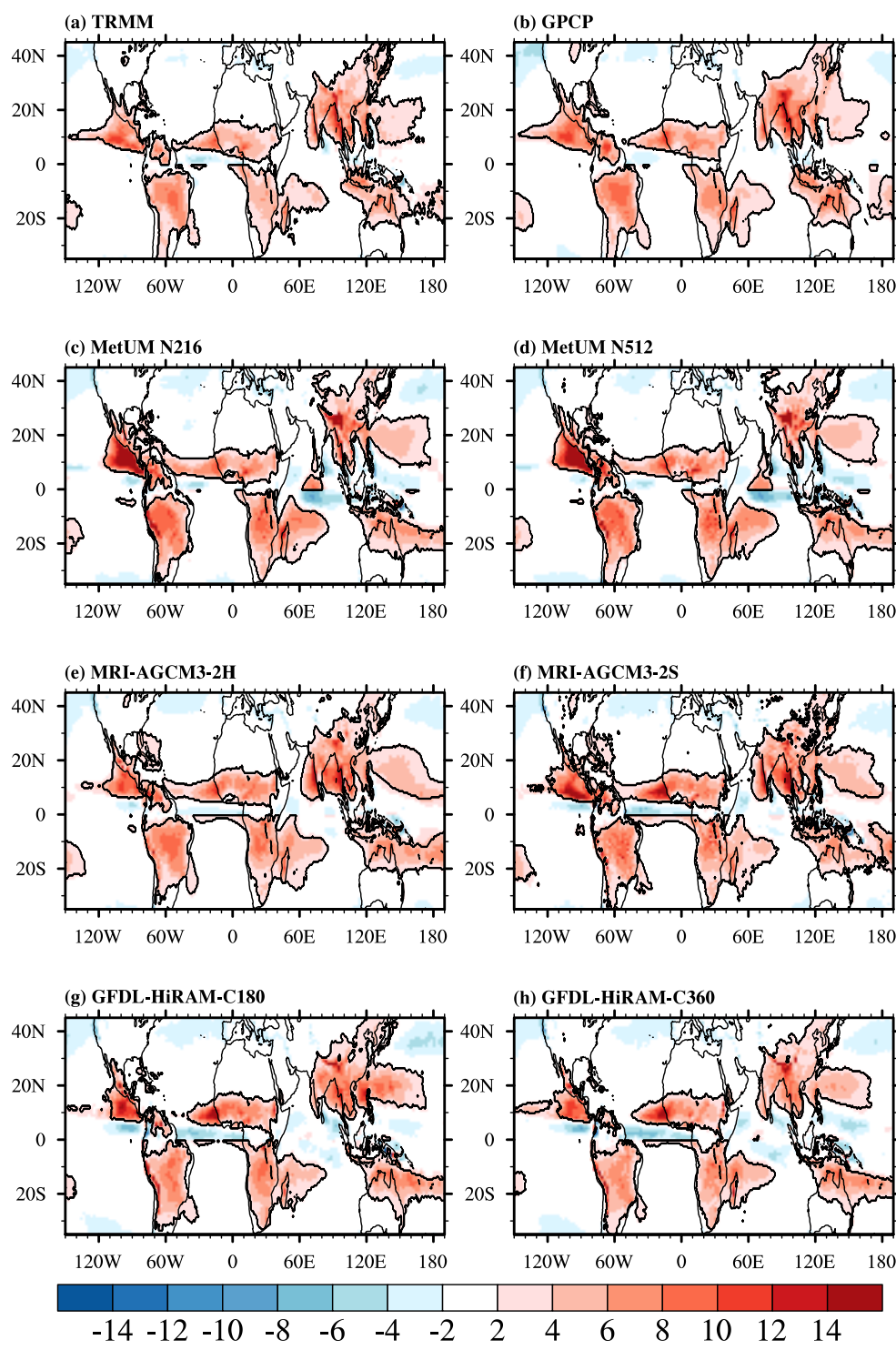


Figure 5 Climatological annual range of precipitation (shaded, units: mm day⁻¹) and global monsoon domain (contours) derived from (a) TRMM, (b) GPCP, (c) MetUM-GA3 N216, (d) MetUM-GA3 N512, (e) MRI-AGCM3-2H, (f) MRI-AGCM3-2S, (g) GFDL-HiRAM-C180, (h) GFDL-HiRAM-C360. The annual range is defined as the May to September (MJJAS) minus November to March (NDJFM) mean precipitation for the Northern Hemisphere, and NDJFM minus MJJAS for the Southern Hemisphere.

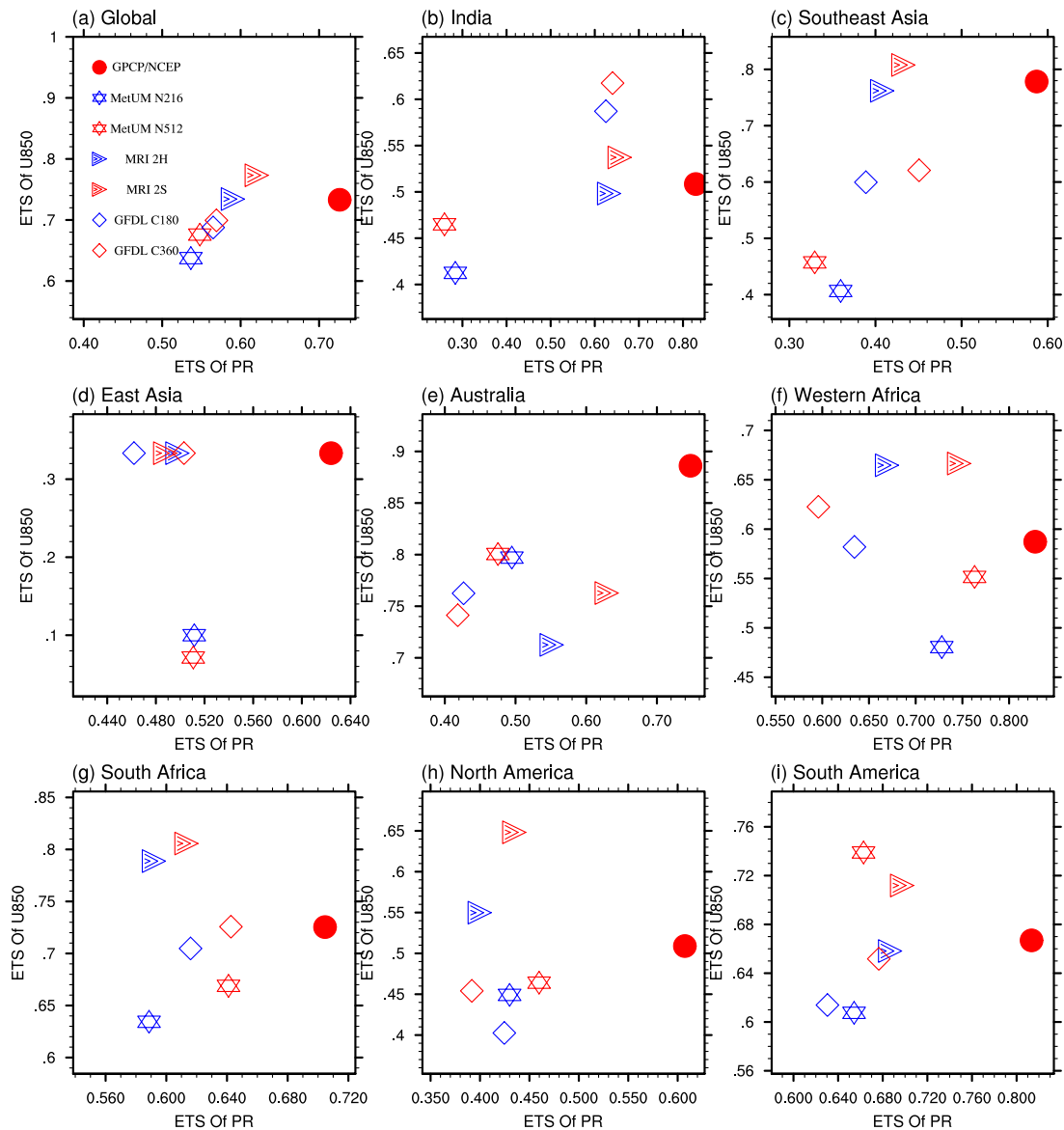


Figure 6. The ETS score of monsoon domain simulation over (a) global (GM), (b) Indian (India), (c) Southeast Asian (SAs), (d) East Asian (EAs), (e) Australian (Aus), (f) West African (Waf), (g) South African (SAf), (h) North American (NAf) and (i) South American (SAf) monsoon. The abscissa (ordinate) is ETS scores of simulated domain derived from precipitation (U850). The fidelity of GPCP against TRMM is shown in dots. The fidelity of GPCP and NCEP2 against TRMM and CFSR are also shown in dots to represent the observational uncertainty. The star, triangle and diamond represent the results from MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution configurations in red and low-resolution configurations in blue.

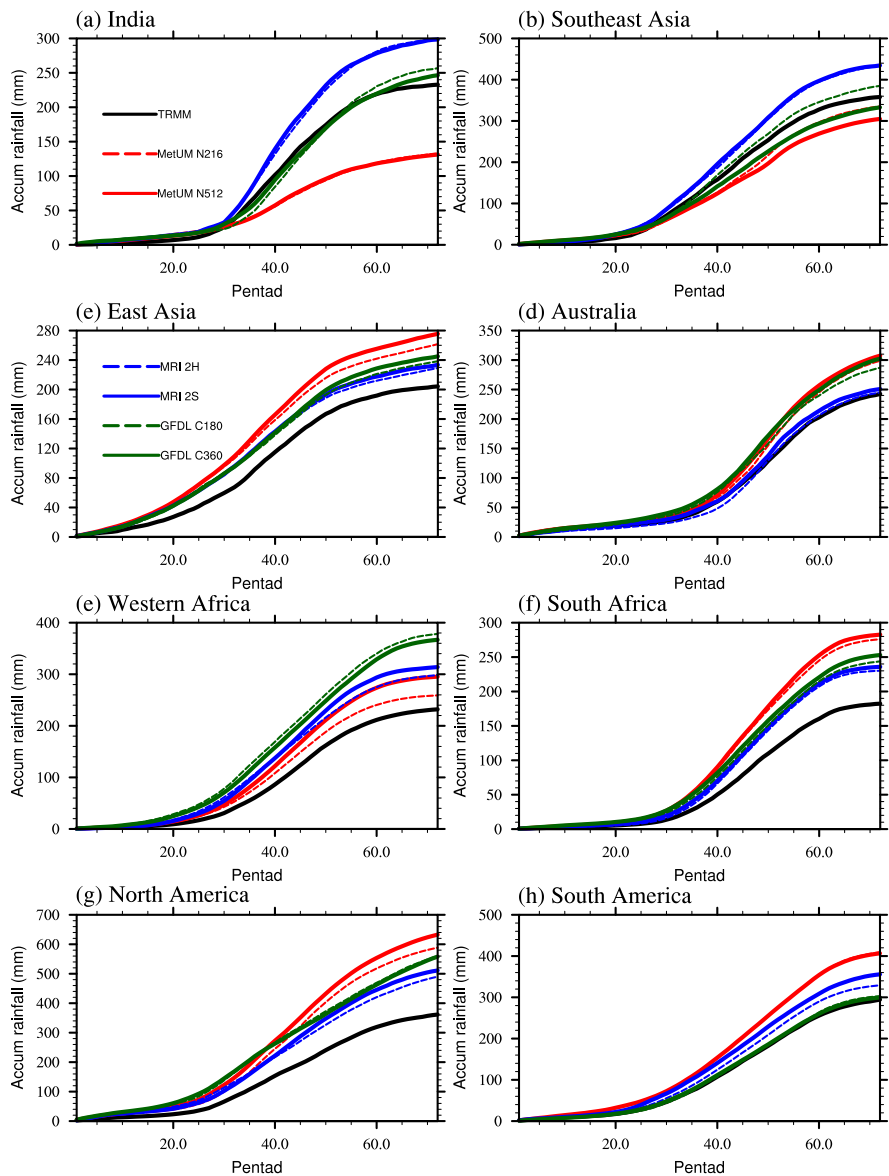


Figure 7 Observed and simulated precipitation accumulations (unit: mm) from the first pentad to the end of year, and fractional accumulations (unit: %) around the monsoon onset pentad averaged over (a) Indian (India), (b) Southeast Asian (SAs), (c) East Asian (EAs), (d) Australian (Aus), (e) West African (Waf), (f) South African (SAf), (g) North American (NA_m) and (h) South American (SA_m) monsoon. Note: for Southern Hemisphere locations (Australian, South African, and South American monsoons) the pentads were reordered to July–June prior to the analysis. The black, red, blue and green lines represent the results of TRMM, MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution configurations in solid lines and low-resolution configurations in dashed lines.

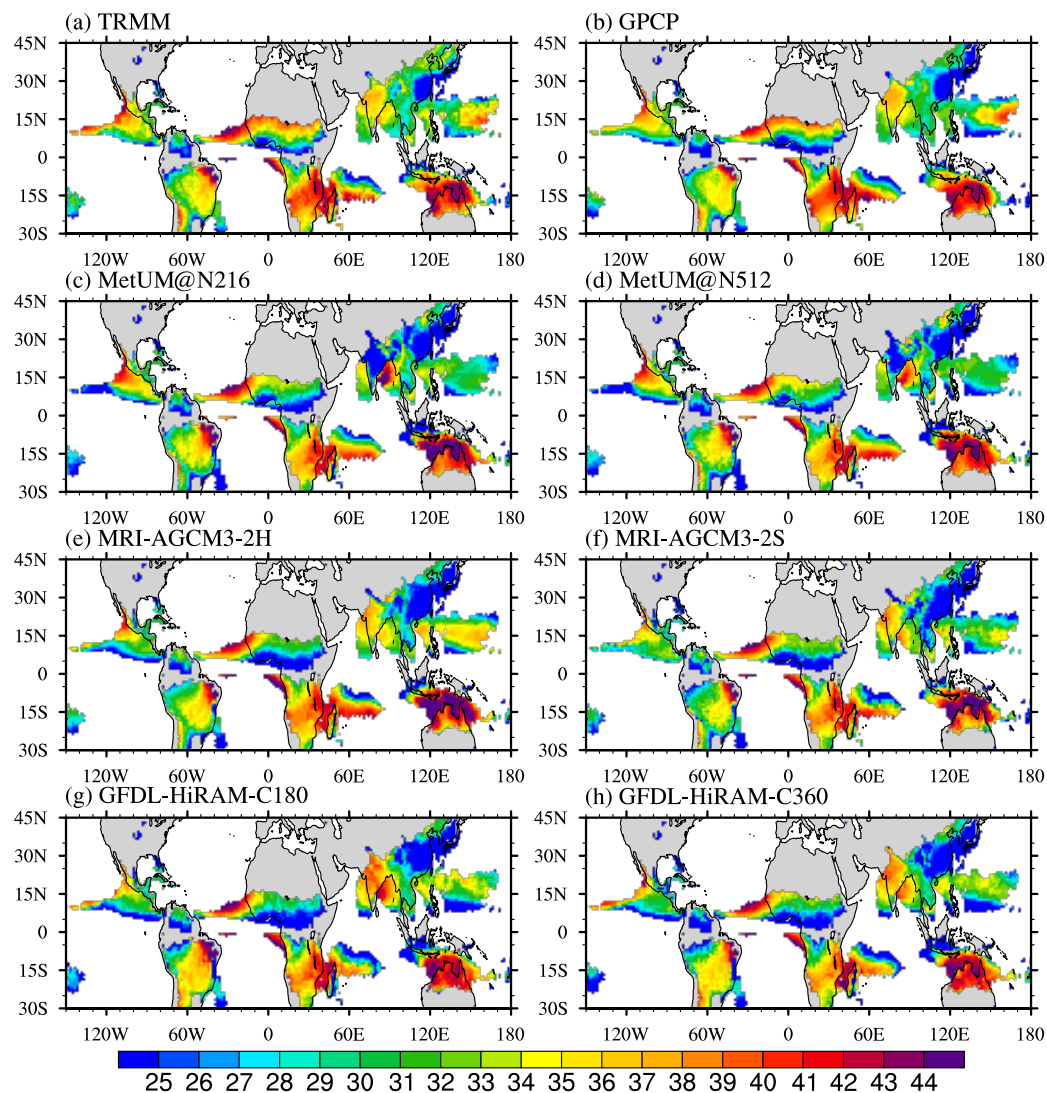


Figure 8 Climatological monsoon onset pentad derived from (a) TRMM, (b) GPCP, (c) MetUM-GA3 N216, (d) MetUM-GA3 N512, (e) MRI-AGCM3-2H, (f) MRI-AGCM3-2S, (g) GFDL-HiRAM-C180, (h) GFDL-HiRAM-C360. The monsoon onset pentad is defined as the pentad when fractional accumulation first becomes ≥ 0.2 (Sperber and Annamalai 2014). Note: for Southern Hemisphere locations (Australian, South African, and South American monsoon) the pentads were reordered to July–June prior to the analysis.

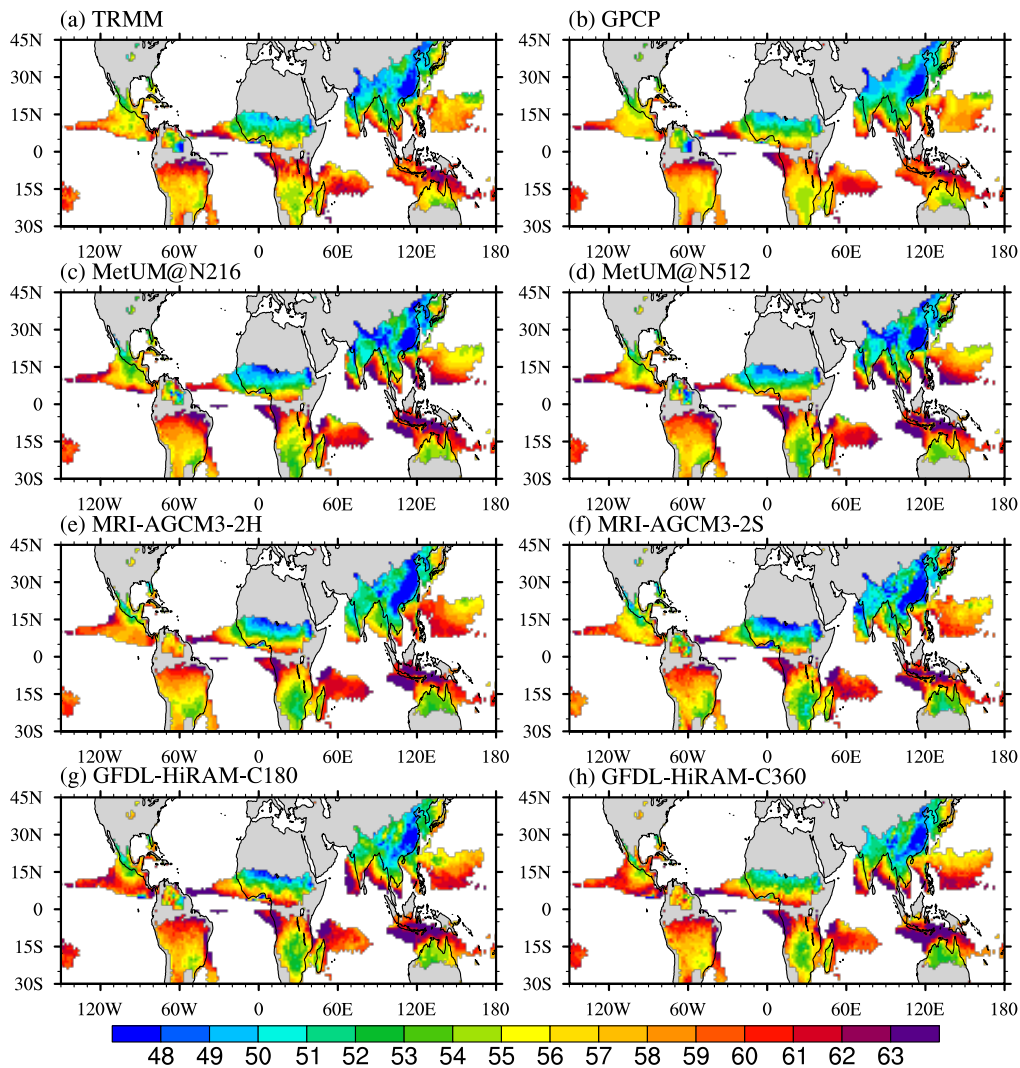


Figure 9 As in Figure 8, but for the monsoon withdrawal pentad, defined as the pentad with fractional accumulation first becomes ≥ 0.8 (Sperber and Annamalai 2014). Note: for Southern Hemisphere locations (Australian, South African, and South American monsoon) the pentads were reordered to July–June prior to the analysis.

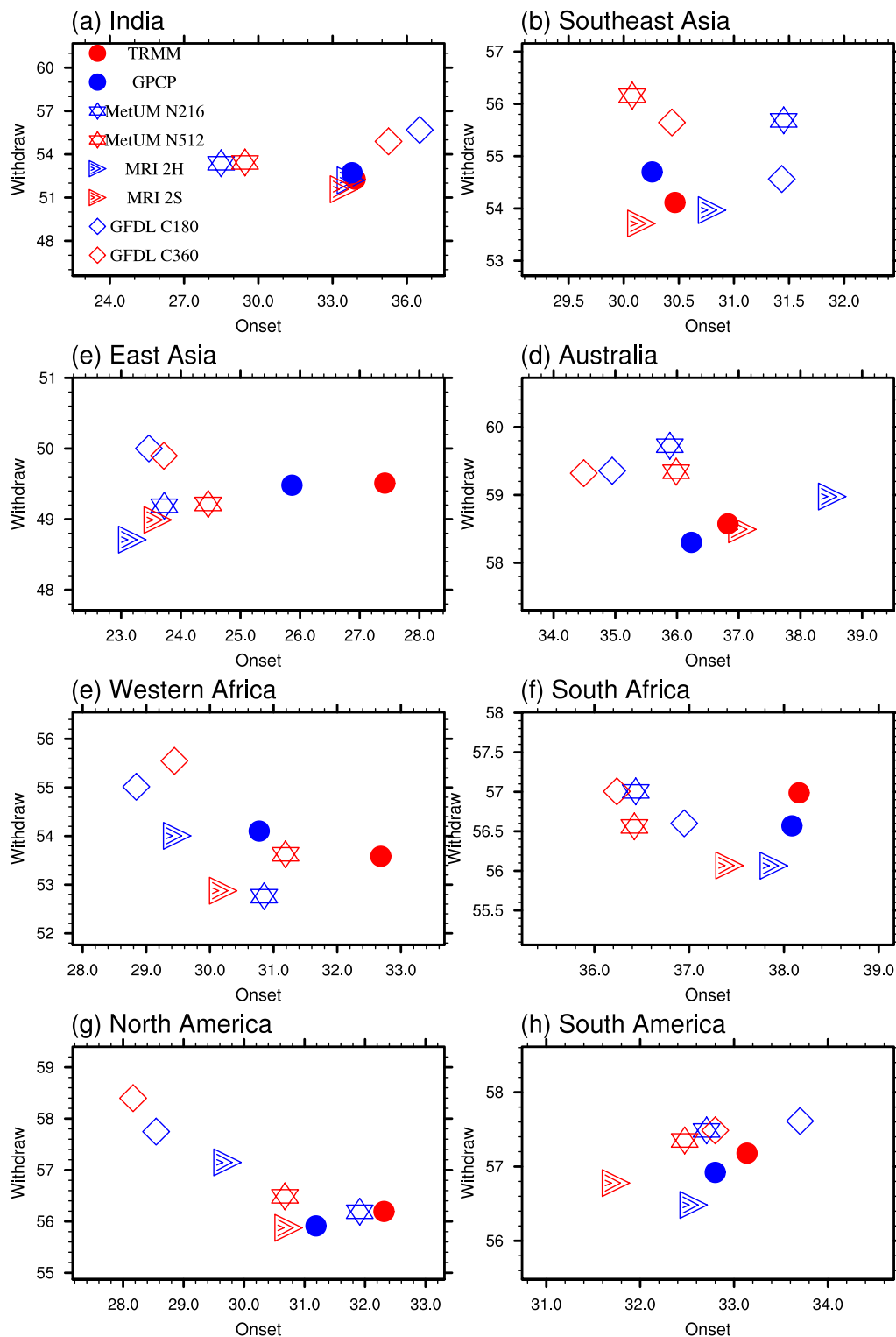


Figure 10. The monsoon onset (x axis) and withdraw pentad (y axis) averaged over (a) Indian (India), (b) Southeast Asian (SAs), (c) East Asian (EAs), (d) Australian (Aus), (e) West African (WAf), (f) South African (SAf), (g) North American (NAf) and (h) South American (SAm) monsoon. Note: for Southern Hemisphere locations (Australian, South

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African, and South American monsoon) the pentads were reordered to July–June prior to the analysis. The dots, stars, triangles and diamonds represent the results from the observations, MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively, with high-resolution configurations in red and low-resolution configurations in blue.

For Review Only

SAm	0.41,0.20,0.25	0.46,0.19,0.30	0.30,0.02,0.19	0.01,0.01,0.04	-0.1,0.1,1.9	-0.1,0.3,-0.1
NAm	0.32,0.16,0.21	0.31,-0.32,0.69	0.10,-0.25,0.44	0.03,0.03,-0.03	-0.3,-1.3,0.7	0.3,-1.3,0.7
SAf	0.35,0.07,0.26	0.24,0.03,0.27	0.10,0.05,0.01	0.04,0.08,-0.04	0.0,-0.5,-0.7	-0.4,0.0,0.4
WAf	0.16,-0.04,0.14	0.25,0.05,0.19	0.15,0.13,0.18	-0.02,0.04,-0.02	0.3,0.7,0.6	0.9,-1.1,0.5
Aus	0.14,-0.03,0.09	0.57,-0.15,0.46	0.32,-0.07,0.16	0.0,-0.01,0.04	0.1,-1.5,-0.5	-0.4,0.0,0.9
EAs	0.21,0.10,0.13	0.50,0.31,0.47	0.05,0.07,0.14	-0.03,0.02,0.06	0.7,0.4,0.3	0.0,0.3,-0.1
SAs	0.34,0.47,0.30	0.12,0.36,0.68	0.18,0.05,0.17	0.02,-0.03,0.02	-1.4,-0.6,-1.0	0.5,-0.3,1.1
India	0.16,0.64,0.52	0.15,0.85,0.72	0.20,0.07,0.25	-0.02,0.02,0.02	1.0,-0.3,-1.3	0.1,-0.6,-0.8
GM	0.25,0.20,0.16	0.09,0.11,0.11	0.15,0.05,0.17	0.01,0.03,0.01	0.0,-0.6,-0.4	0.2,-0.5,0.1
	ANN	AC1	AC2	Domain	Onset	WithD

Figure 11 Improvement of the simulation on the annual cycle of global monsoon precipitation with horizontal resolution. Red boxes indicate high-resolution consistently improves performance across all three AGCMs; blue boxes indicate improvements with resolution are inconsistent; orange boxes indicate a better simulation is shown in all three high-resolution AGCMs but the responses to resolution are inconsistent across models. The values from left to right in each box are the difference between high- and low-resolution configurations of MetUM-GA3, MRI-AGCM3 and GFDL-HiRAM, respectively. The metric for the annual modes, including annual mean (ANN), solstitial mode (AC1), the equinoctial asymmetric mode (AC2), is the pattern correlation coefficient with TRMM. The metrics for the monsoon domain and monsoon onset/withdrawal are ETS and the onset/withdrawal pentad, respectively.